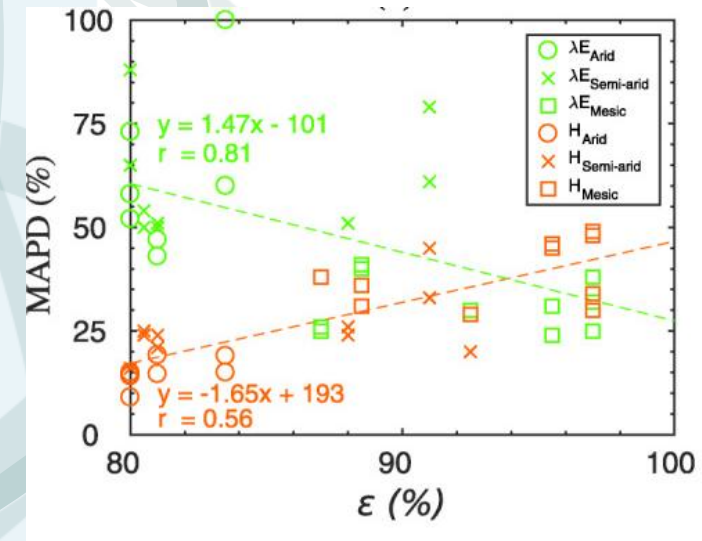


Evapotranspiration mapping across an aridity gradient in conterminous US by combining thermal remote sensing with Penman-Monteith and Shuttleworth-Wallace model

Kanis(h)ka Mallick¹, Nishan Bhattarai²,
Nathaniel Brunsel³, Ge Sun⁴, Meha Jain²



¹Water Safety and Security Research Unit, Department ERIN, Luxembourg Institute of Science and Technology (LIST)

²School for Environment and Sustainability, University of Michigan, USA

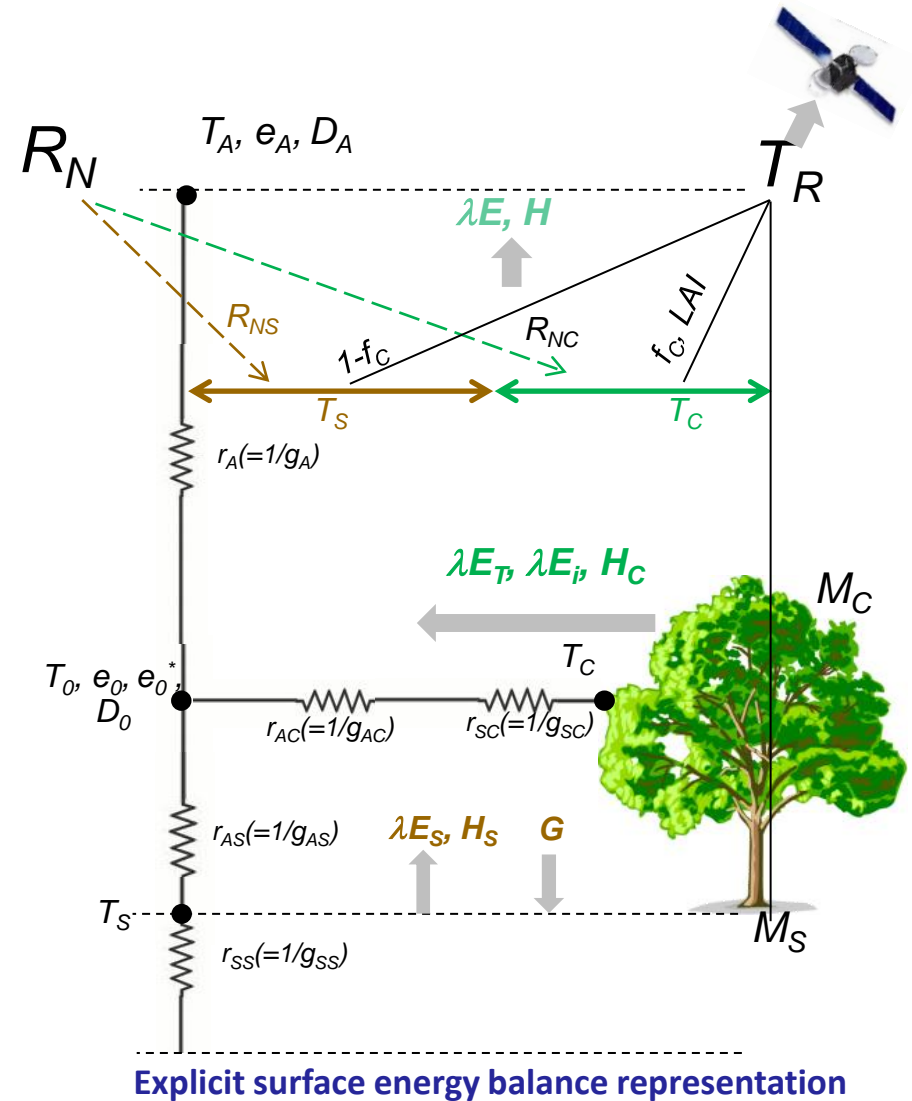
³Geography and Atmospheric Science, University of Kansas, USA

⁴Eastern Forest Environmental Threat Assessment Center, Southern Research Station, US Department of Agriculture Forest Service, Raleigh, USA

kaniska.mallick@gmail.com;
kaniska.mallick@list.lu

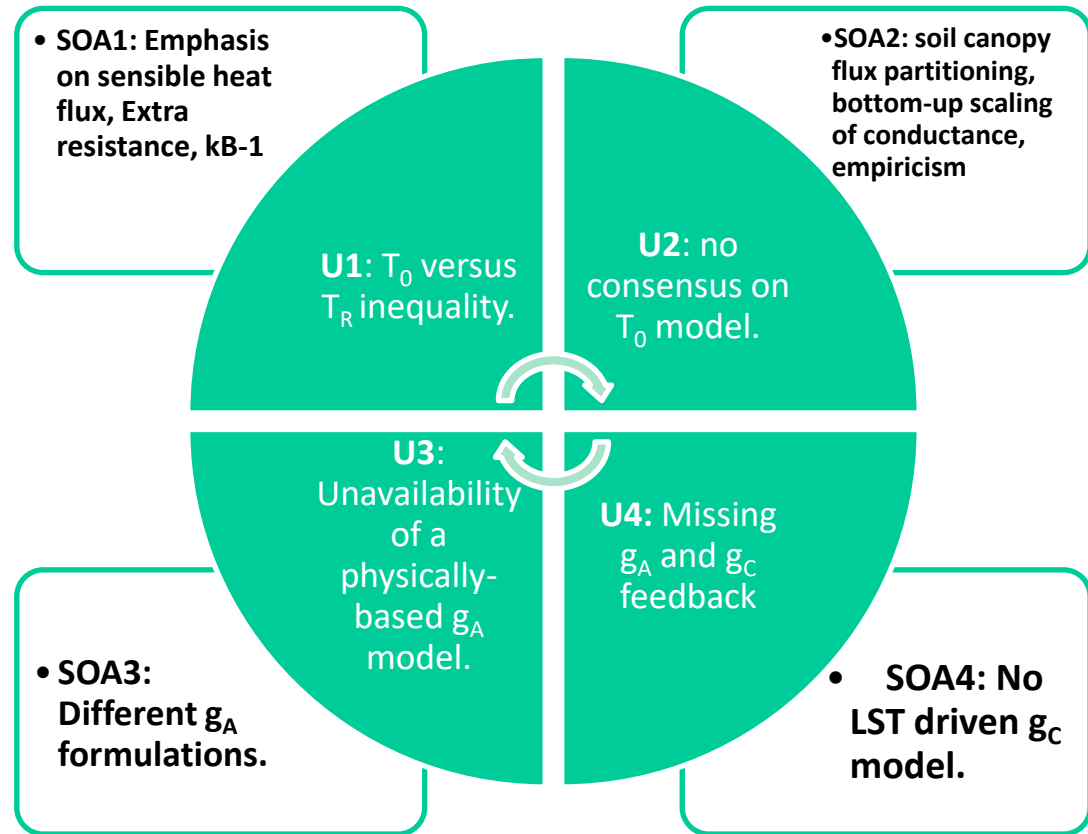
Summary

- Challenges in thermal remote sensing of ET.
- Why Penman-Monteith (PM) and Shuttleworth-Wallace (SW)?
- Proposed modeling scheme and characteristics
- Study region and data
- Results
- Conclusion



State-of-the-art uncertainties / challenges

- Inequality between aerodynamic temperature (T_0) and T_R ($T_0 \neq T_R$)
- Non-unique relationship between T_0 and T_R
- universally agreed T_0 model: unavailable
- Aerodynamic conductance (g_A):
Semi-empirical
- Canopy conductance (g_C):
oversighting the role of LST on g_C .

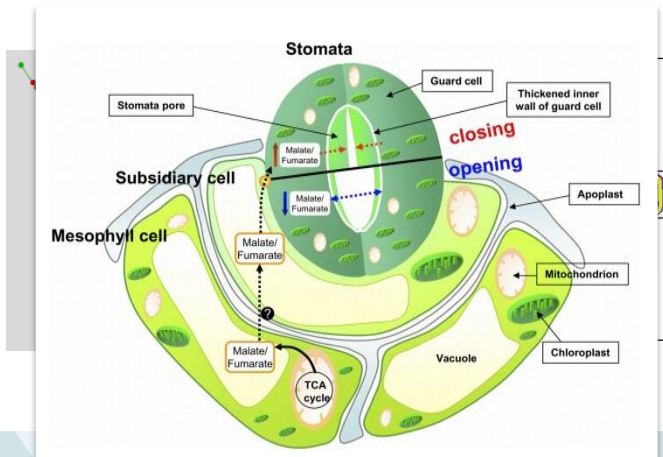


Why PM and SW?

Penman-Monteith (Monteith, 1965, 1981)

$$\lambda \dot{E}_{PM} = \frac{s\phi + \rho c_P g_A D_A}{s + \gamma \left(1 + \frac{g_A}{g_c}\right)}$$

$$\phi = R_N - G, s = f\{T_A\}$$



Shuttleworth-Wallace (Shuttleworth and Wallace, 1985)

$$\lambda \dot{E}_{SW} = \frac{s\phi_C + \rho c_P g_A^C D_0}{s + \gamma \left(1 + \frac{g_A^C}{g_S^C}\right)} + \frac{s\phi_S + \rho c_P g_A^S D_0}{s + \gamma \left(1 + \frac{g_A^S}{g_S^S}\right)}$$

$$D_0 = D_A + \frac{\{s\phi - (s + \gamma)\lambda \dot{E}_{PM}\} g_A}{\rho c_P}$$

g_A = aerodynamic conductance
 g_c = canopy (surface) conductance

Integrating LST into PM-SW

STIC (Surface Temperature Initiated Closure)

$$\lambda E_{PM} = \frac{s\phi + \rho c_P g_A D_A}{s + \gamma \left(1 + \frac{g_A}{g_C}\right)}$$

$$sv_1 = f\{c_1, c_2, c_3, v_1, v_2, v_3, v_4, sv_3, sv_5\}$$

$$sv_2 = f\{v_3, v_4, sv_1, sv_5, sv_6\}$$

$$sv_3 = f\{c_3, v_3, v_4, sv_4, sv_5\}$$

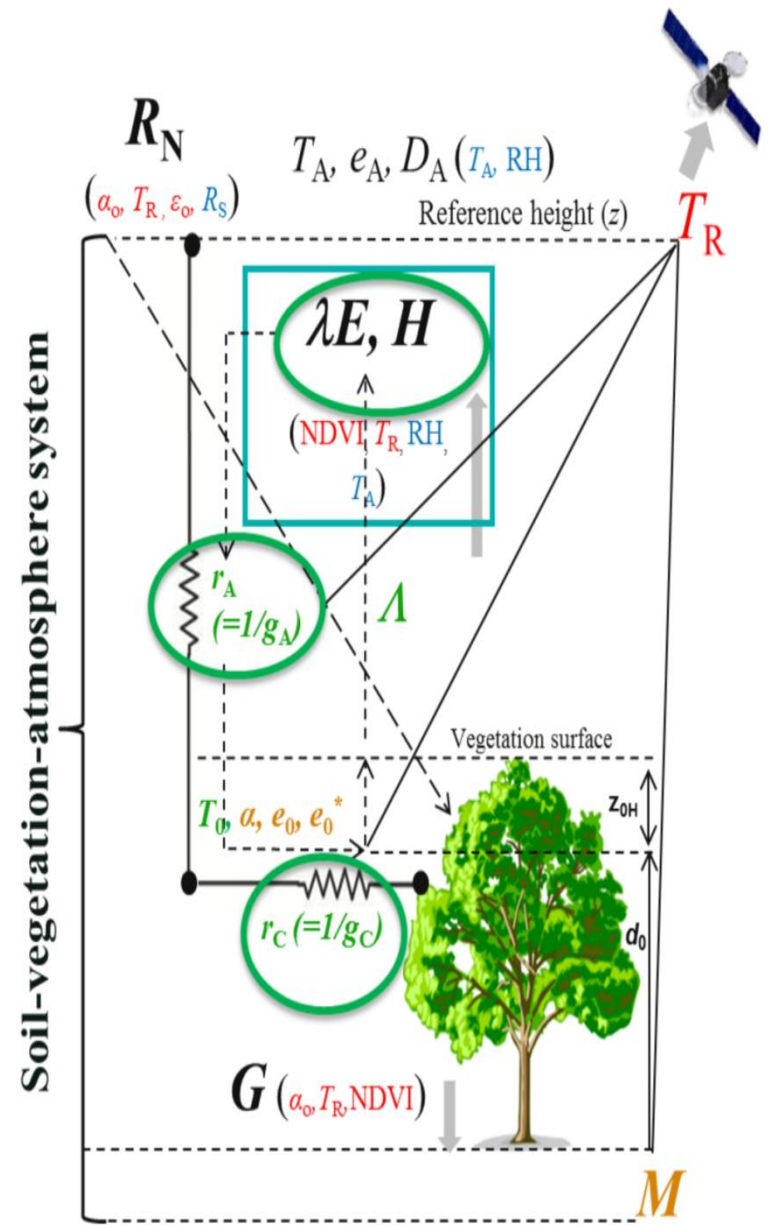
$$sv_4 = f\{c_3, v_3, sv_1, sv_2, sv_7, sv_8\}$$

$$sv_5 = f\{c_1, c_2, c_3, v_1, v_2, v_3, v_4, sv_1, sv_2, \lambda E_{PM}\}$$

$$sv_6 = f\{c_1, c_2, c_3, v_3, v_4, sv_1, sv_2, \lambda E_{PM}\}$$

$$sv_7 = f\{c_1, c_2, c_3, v_3, v_4, sv_1, sv_2, sv_3, sv_6, sv_8\}$$

$$sv_8 = f\{v_3, v_4, v_5, sv_1, \lambda E_{PM}\}$$



1-D Surface energy balance representation

Bhattarai et al., 2018; Mallick et al., 2016, 2018

Characteristics

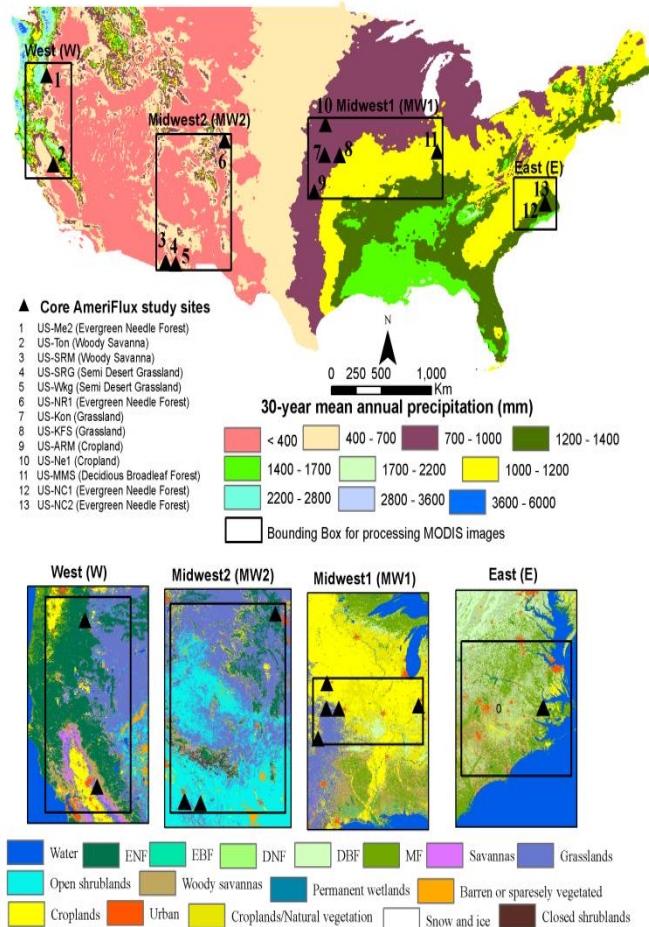
Functional

- Fully analytical
- LST, aerodynamic conductance and vapor pressure feedback
- Simultaneous ET partitioning
- Application potential: both LEO and GEO

Structural

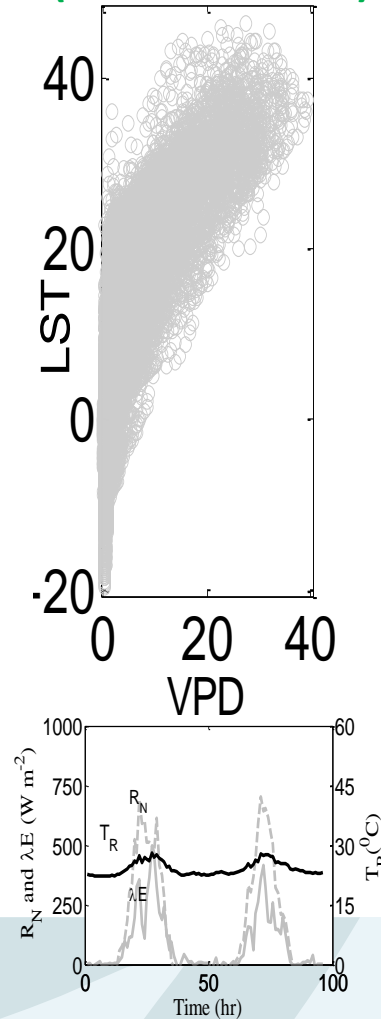
- Physical integration of LST: combining PM and SW to solve D_0
- No land surface parameterization for the conductances
- Direct estimation of ET and H
- Numerical estimation: Conductances, Priestley-Taylor α (as a time varying quantity, instead of a fixed value), canopy-air stream properties.
- Inputs: R_N , G , T_A , R_H or e_A , and LST (or TR).

Evaluation: across an aridity gradient ([Bhattarai, Mallick et al., 2018](#))

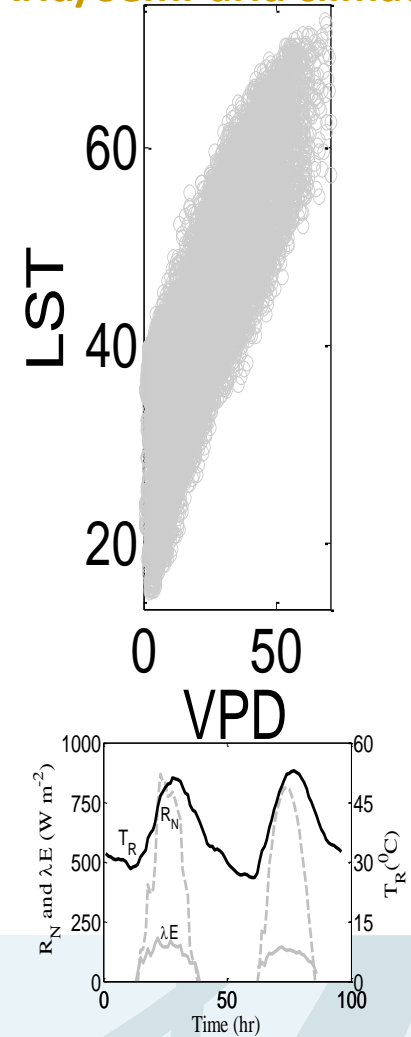


Data : MODIS LST, surface reflectances, NLDAS meteorology

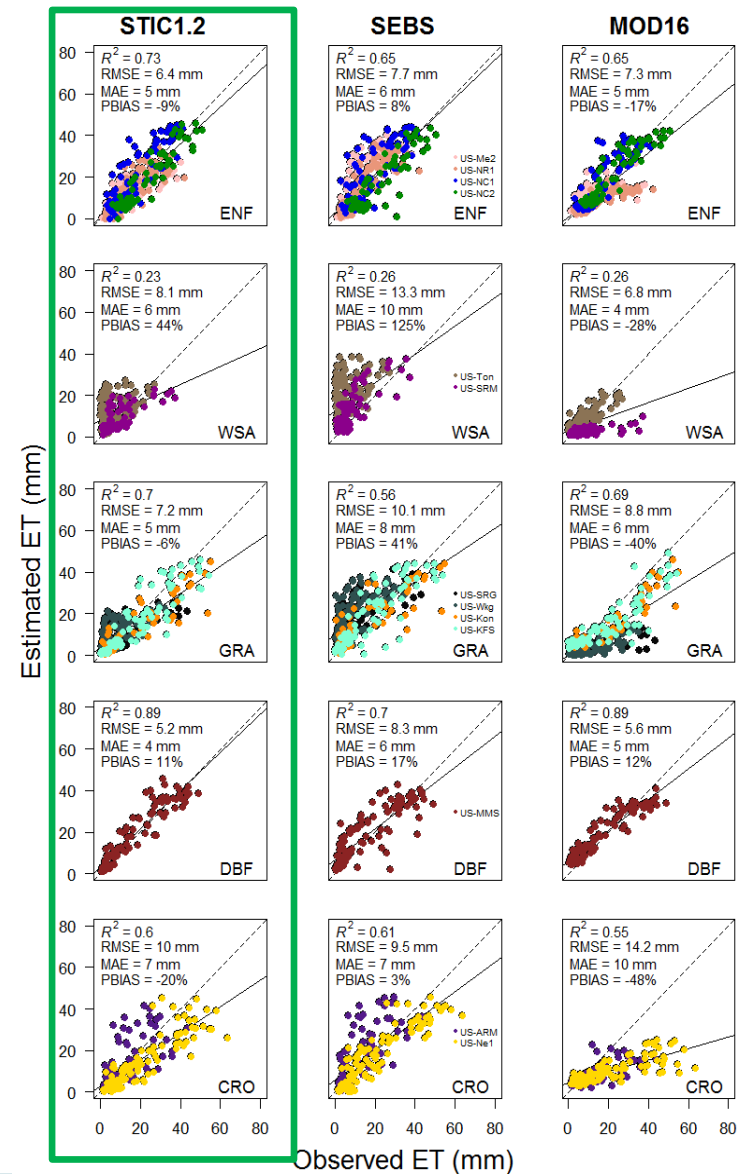
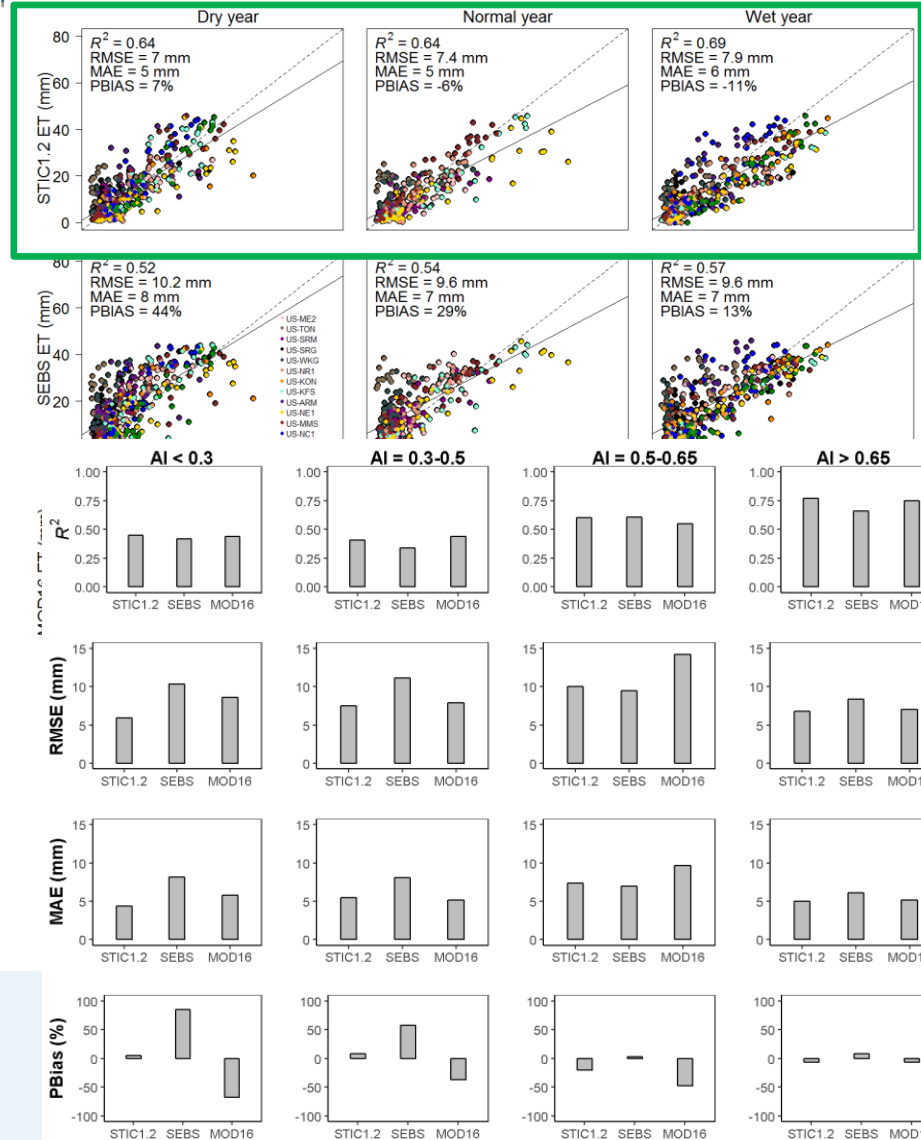
**Radiation controlled
(Humid climate)**



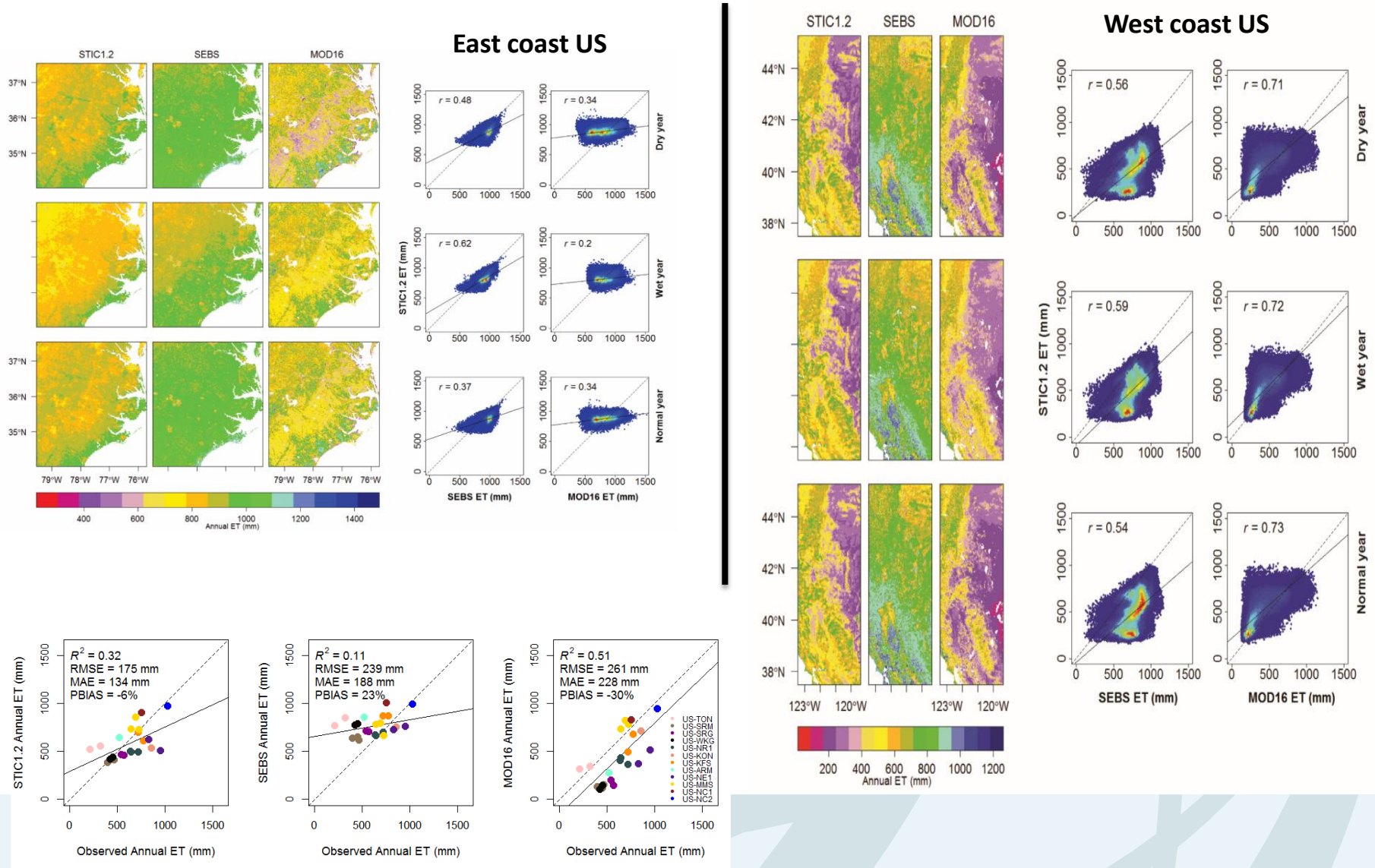
**Water controlled
(Arid/Semi-arid climate)**



Evaluation and model intercomparison (Precipitation extreme, biome, aridity)

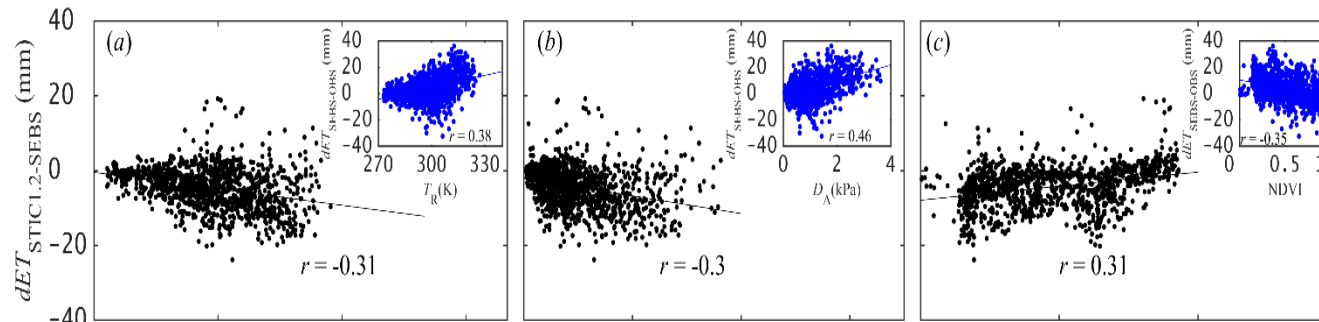


Annual ET distribution and evaluation

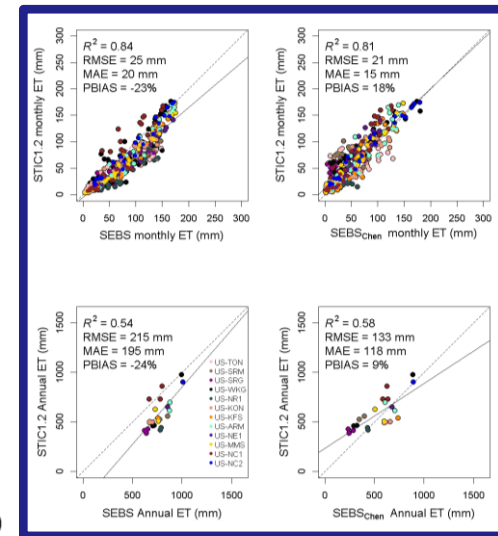
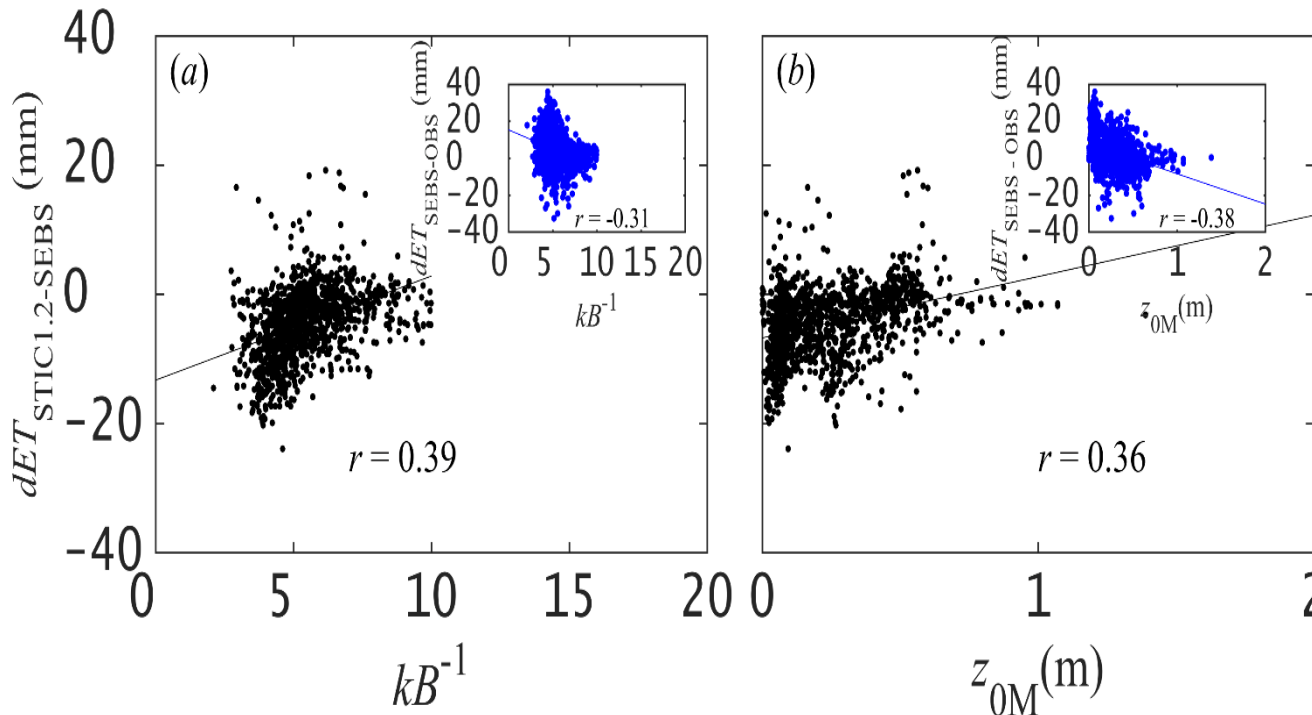


Model differences: Forcings versus parameterizations

(Bhattarai, Mallick et al., 2018)



- Overestimation SEBS: high T_R and D_A , low NDVI.
- Major ET difference: kB^{-1} (2 – 6)



Interpretation

- STIC1.2 explained significant variability in the observed 8-day cumulative ET, RMSE<1 mm/d
- Smallest errors in forests, followed by grassland, cropland, and woody savannas.
- Underestimation of ET in croplands: spatial-scale mismatch between a MODIS pixel and the flux tower footprint
- Overestimation of ET in woody savannas: large uncertainties in the MODIS LST product, SEB closure correction of EC ET observations, single-source approximations.
- Difference between STIC1.2 and SEBS: Differences in g_A estimation between the two models.
- Empirical characterization of z_{0M} and kB^{-1} in SEBS: major factors creating uncertainties in aerodynamic conductance and ET estimations.

Thank you!!

Hydrol. Earth Syst. Sci., 22, 2311–2341, 2018
https://doi.org/10.5194/hess-22-2311-2018
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Regional evapotranspiration from an image-based implementation of the Surface Temperature Initiated Closure (STIC1.2) model and its validation across an aridity gradient in the conterminous US

Nishan Bhattarai¹, Kaniska Mallick², Nathaniel A. Brunsell³, Ge Sun⁴, and Meha Jain¹

¹School for Environment and Sustainability, University of Michigan, Ann Arbor, MI 48109, USA

²Remote Sensing and Ecophysiological Modeling, Water Security and Safety Research Unit, Dept. ERIN, Luxembourg Institute of Science and Technology (LIST), 4422 Belvaux, Luxembourg

³Geography and Atmospheric Science, University of Kansas, Lawrence, KS 66045, USA

⁴Eastern Forest Environmental Threat Assessment Center, Southern Research Station, US Department of Agriculture Forest Service, Raleigh, NC 27606, USA

Correspondence: Nishan Bhattarai (nbhattar@umich.edu)

Received: 30 August 2017 – Discussion started: 11 September 2017

Revised: 19 March 2018 – Accepted: 19 March 2018 – Published: 18 April 2018

Hydrol. Earth Syst. Sci., 20, 4237–4264, 2016
www.hydrol-earth-syst-sci.net/20/4237/2016/
doi:10.5194/hess-20-4237-2016

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Canopy-scale biophysical controls of transpiration and evaporation in the Amazon Basin

Kaniska Mallick¹, Ivonne Trebs¹, Eva Boegh², Laura Glustarini¹, Martin Schlerf¹, Darren T. Drewry^{3,12}, Lucien Hoffmann⁴, Celso von Randow⁴, Bart Kruijt⁴, Alessandro Araújo⁴, Scott Saleska⁵, James R. Ehleringer⁶, Tomas F. Domingues⁷, Jean Pierre H. B. Ometto⁸, Antonio D. Nobre⁴, Osvaldo Luiz Leal de Moraes¹⁰, Matthew Hayek¹¹, J. William Munger¹¹, and Steven C. Wofsy¹¹

¹Department of Environmental Research and Innovation, Luxembourg Institute of Science and Technology (LIST), L4422, Belvaux, Luxembourg

²Department of Science and Environment, Roskilde University, Roskilde, Denmark

³Jet Propulsion Laboratory, California Institute of Technology, 4800 Oak Grove Drive, Pasadena, 91109, USA

⁴Instituto Nacional de Pesquisas Espaciais (INPE), Centro de Ciência do Sistema Terrestre, São José dos Campos, SP, Brazil

⁵Wageningen Environmental Research (ALTEIRA), Wageningen, the Netherlands

⁶Empresa Brasileira de Pesquisa Agropecuária (EMBRAPA), Belém, PA, Brazil

⁷Department of Ecology and Evolutionary Biology, University of Arizona, Tucson, AZ, USA

⁸Department of Biology, University of Utah, Salt Lake City, UT, USA

⁹Faculdade de Filosofia Ciências e Letras de Ribeirão Preto, Universidade de São Paulo (USP), São Paulo, SP, Brazil

¹⁰Centro Nacional de Monitoramento e Alertas de Desastres Naturais, São Paulo, SP, Brazil

¹¹Department of Earth and Planetary Science, Harvard University, Cambridge, MA, USA

¹²Joint Institute for Regional Earth System Science and Engineering, University of California, Los Angeles, California, USA

Correspondence to: Kaniska Mallick (kaniska.mallick@gmail.com) and Ivonne Trebs (ivonne.trebs@list.lu)

Received: 30 December 2015 – Published in Hydrol. Earth Syst. Sci. Discuss.: 27 January 2016

Revised: 21 June 2016 – Accepted: 14 September 2016 – Published: 19 October 2016

Hydrology and
Earth System
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AGU100 ADVANCING
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Water Resources Research

RESEARCH ARTICLE

10.1029/2017WR021357

Special Section:

Hydrology Delivers Earth
System Sciences to Society
(HESSS4): Improving and
Integrating Knowledge Across
Disciplines on Global Energy,
Water and Carbon Cycles

Key Points:

- Thermal remote sensing of evapotranspiration is critical due to uncertainties in aerodynamic temperature and conductance estimation
- We integrated radiometric temperature into Penman-Monteith Shuttleworth-Wallace framework to directly estimate conductances and evapotranspiration
- Moderate to low systematic errors in evapotranspiration across an aridity gradient in Australia

Bridging Thermal Infrared Sensing and Physically-Based Evapotranspiration Modeling: From Theoretical Implementation to Validation Across an Aridity Gradient in Australian Ecosystems

Kaniska Mallick¹, Erika Toivonen^{1,2,3,4}, Ivonne Trebs¹, Eva Boegh^{5,6}, James Cleverly⁷, Derek Eamus⁷, Harri Koivusalo², Darren Drewry^{8,9}, Stefan K. Arndt¹⁰, Anne Griebel¹⁰, Jason Beringer¹¹, and Monica Garcia^{12,13}

¹Department of Environmental Research and Innovation, Luxembourg Institute of Science and Technology, Belvaux, Luxembourg, ²Department of Built Environment, Aalto University School of Engineering, Espoo, Finland, ³Climate System Research, Finnish Meteorological Institute, Helsinki, Finland, ⁴Department of Physics, University of Helsinki, Helsinki, Finland, ⁵Department of Science and Environment, Roskilde University, Roskilde, Denmark, ⁶Now at Danish Agency for Data Supply and Efficiency, Copenhagen, Denmark, ⁷Terrestrial Ecophysiology Research Group, School of Life Sciences, University of Technology Sydney, Broadway, NSW, Australia, ⁸Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA, USA, ⁹Joint Institute for Regional Earth System Science and Engineering, University of California, Los Angeles, CA, USA, ¹⁰School of Ecosystem and Forest Sciences, University of Melbourne, Melbourne, Vic, Australia, ¹¹School of Agriculture and Environment, University of Western Australia, Crawley, WA, Australia, ¹²Department of Environmental Engineering, Technical University of Denmark, Kgs. Lyngby, Denmark, ¹³International Research Institute for Climate and Society, Earth Institute, Columbia University, Palisades, NY, USA

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Water Resources Research

RESEARCH ARTICLE

10.1002/2014WR016106

Key Points:

- Reintroducing radiometric surface temperature into Penman-Monteith (PM) model
- Holistic surface moisture availability framework to constrain the PM equation
- Numerical estimation of Priestley-Taylor parameter

Correspondence to:

K. Mallick,
kaniska.mallick@gmail.com;
kaniska.mallick@list.lu

Citation:
Mallick, K., E. Boegh, I. Trebs,
J. G. Alfieri, W. P. Roushi, J. H. Prueger,
D. Niyogi, N. Das, D. T. Drewry,
L. Hoffmann, and A. J. Jarvis (2015),
Reintroducing radiometric surface
temperature into the Penman-
Monteith formulation, *Water Resour.
Res.*, 51, 6214–6243, doi:10.1002/
2014WR016106.

Received 7 JUL 2014
Accepted 9 JUL 2015
Accepted article online 14 JUL 2015
Published online 8 AUG 2015
Corrected 27 AUG 2015

Reintroducing radiometric surface temperature into the Penman-Monteith formulation

Kaniska Mallick¹, Eva Boegh², Ivonne Trebs¹, Joseph G. Alfieri³, William P. Roushi⁴, John H. Prueger⁵, Dev Niyogi⁶, Narendra Das⁶, Darren T. Drewry⁶, Lucien Hoffmann⁷, and Andrew J. Jarvis⁸

¹Department of Environmental Research and Innovation, Luxembourg Institute of Science and Technology, Belvaux, Luxembourg, ²Department of Environmental, Social and Spatial Change, Roskilde University, Roskilde, Denmark, ³USDA-ARS, Hydrology and Remote Sensing Laboratory, Beltsville, Maryland, USA, ⁴USDA-ARS, National Laboratory for Agriculture and Environment, Ames, Iowa, USA, ⁵Department of Agronomy and the Department of Earth and Atmospheric and Planetary Sciences, Purdue University, West Lafayette, Indiana, USA, ⁶Jet Propulsion Laboratory, California Institute of Technology, Pasadena, California, USA, ⁷Lancaster Environment Centre, Lancaster University, Lancaster, UK

Abstract Here we demonstrate a novel method to physically integrate radiometric surface temperature (T_a) into the Penman-Monteith (PM) formulation for estimating the terrestrial sensible and latent heat fluxes (H and LE) in the framework of a modified Surface Temperature Initiated Closure (STIC). It combines T_a data with standard energy balance closure models for deriving a hybrid scheme that does not require parameterization of the surface (or stomatal) and aerodynamic conductances (g_s and g_a). STIC is formed by the simultaneous solution of four state equations and it uses T_a as an additional data source for retrieving the “near surface” moisture availability (W) and the Priestley-Taylor coefficient (s). The performance of STIC is tested using high-temporal resolution T_a observations collected from different international surface energy flux experiments in conjunction with corresponding net radiation (R_n), ground heat flux (G), air temperature (T_a), and relative humidity (R_h) measurements. A comparison of the STIC outputs with the eddy covariance measurements of LE and H revealed RMSDs of 7–16% and 40–74% in half-hourly LE and H estimates. These statistics were 5–13% and 10–44% in daily LE and H . The errors and uncertainties in both surface fluxes are comparable to the models that typically use land surface parameterizations for determining the unobserved components (g_s and g_a) of the surface energy balance models. However, the scheme is simpler, has the capabilities for generating spatially explicit surface energy fluxes and independent of submodels for boundary layer developments.